On the precision of noise correlation interferometry

Richard L. Weaver, 1 Céline Hadziioannou, 2 Eric Larose 2 and Michel Campillo 2

1 Department of Physics, University of Illinois at Urbana-Champaign, 1110 W Green, Urbana, Illinois 61801, USA
2 Institut des Sciences de la Terre (ISTerre), Université J. Fourier & CNRS, BP 53, 38041 Grenoble cedex 9, France. E-mail: eric.larose@ujf-grenoble.fr

Accepted 2011 March 10. Received 2011 February 10; in original form 2010 November 29

SUMMARY
Long duration noisy-looking waveforms such as those obtained in randomly multiple scattering and reverberant media are complex; they resist direct interpretation. Nevertheless, such waveforms are sensitive to small changes in the source of the waves or in the medium in which they propagate. Monitoring such waveforms, whether obtained directly or obtained indirectly by noise correlation, is emerging as a technique for detecting changes in media. Interpretation of changes is in principle problematic; it is not always clear whether a change is due to sources or to the medium. Of particular interest is the detection of small changes in propagation speeds. An expression is derived here for the apparent, but illusory, waveform dilation due to a change of source. The expression permits changes in waveforms due to changes in wave speed to be distinguished with high precision from changes due to other reasons. The theory is successfully compared with analysis of a laboratory ultrasonic data set and a seismic data set from Parkfield California.

Key words: Time series analysis; Coda waves; Theoretical seismology; Wave scattering and diffraction; Rheology: crusts and lithosphere.

1 INTRODUCTION
The technique proposed in the 1980s (Poupinet et al. 1984) and later called ‘Coda wave interferometry’ (Snieder et al. 2002) compares coda waveforms from multiply scattered waves obtained under different circumstances or on different dates and detects changes in a medium. A multiply scattered wave can resist detailed interpretation, but for purposes of monitoring one may not need to interpret the waveform: it is sufficient to note changes. Coda wave interferometry was first suggested for seismic waves but has also been applied in laboratory ultrasonics (Weaver & Lobkis 2000; De Rosny & Roux 2001; Lu & Michaels 2005; Gorin et al. 2006; Lobkis & Weaver 2008). In many such cases the change is due to a uniform change of temperature, and thus a uniform change in wave velocity. To detect such changes, Weaver & Lobkis (2000) constructed a dilation correlation coefficient between waveforms $\phi_1$ and $\phi_2$.

$$X(\epsilon) = \frac{\int \frac{\phi_1(t)\phi_2(t(1+\epsilon))dt}{\sqrt{\int \phi_1^2(t)dt \int \phi_2^2(t(1+\epsilon))dt}}}{1}$$

$X$ takes on a value of unity at $\epsilon = 0$ if the two waveforms are identical. It will reach a value of unity at some characteristic value of $\epsilon$ if the two waveforms differ only by some temporal dilation. The estimated degree of dilation between two waveforms is taken to be the value of $\epsilon$ at which $X$ is maximum. $X$ reaches a maximum of less than unity if the waveforms differ by more than dilation alone. Therefore, the value of $X$ at its maximum, if it is less than unity, may be interpreted as a measure of the distortion between the waveforms.

An alternative formulation is Poupinet’s doublet method (Poupinet et al. 1984), which breaks $\phi_1$ and $\phi_2$ into a series of short time windows at several distinct times $t$ and determines the apparent shift $\delta t$ between them by examining conventional cross spectrum. $\delta t$ as a function of $t$, and in particular its slope $\delta t/\delta t$ reveals a change in the medium. Poupinet developed the doublet method in which seismic signals from repeated seismic events could be compared to infer changes in the earth (Poupinet et al. 1984). Song & Richards (1996) and Zhang et al. (2005) used this to show that certain earth crossing rays were shifted and distorted compared to versions some years earlier, indicating a relative rotation between the earth and its core.

The extensive literature in recent years on correlations of diffuse acoustic noise has reported theory and measurements in support of the notion that such correlations are essentially equal to the acoustic response that one would have at one receiver were there a source at the other (Lobkis & Weaver 2001; Derode et al. 2003; Snieder 2004; Weaver & Lobkis 2004; Roux et al. 2005; Gouéard et al. 2008; Tsai 2009). More technically, what is recovered is the Green’s function as filtered into the frequency band of the noise, whitened and symmetrized in time. Two different kinds of records can be correlated. Sometimes it is coda that is correlated (Campillo & Paul 2003). Coda waves consist in a long duration noisy-looking waveforms such as those obtained in randomly multiple scattering and reverberant media.
scattering. More commonly the diffuse noise is due to ambient seismic waves from continuously acting sources such as human activity or ocean storms. Much recent literature reports constructions of the earth’s seismic response between two seismograph stations, without the use of controlled sources, and without waiting for a seismic event. Tomographic maps of seismic velocity with unprecedented resolution have been obtained (Sabra et al. 2005; Shapiro et al. 2005). The technique has even been applied on the moon (Larose et al. 2005). Very commonly, the noise which is correlated is incompletely equipartitioned, such that the resulting correlation waveforms do not precisely correspond to the Green’s function. A difference between noise correlation and Green’s function can also be observed when one has not averaged enough raw data; the correlation may not have yet converged. Theoretical and applied work is ongoing in attempts to understand and correct for systematic errors due to these effects (Weaver et al. 2009; Froment et al. 2010). Nevertheless, Hadziioannou et al. (2009) demonstrated that it is not necessary to reconstruct the Green’s function to use correlations for monitoring purposes.

These two approaches have been combined into what may be termed noise-correlation interferometry (Sabra et al. 2006; Sens-Schönfelder & Wegler 2006; Brenguier et al. 2008a; Brenguier et al. 2008b) in which correlations of seismic noise taken in different circumstances are compared. The correlations may have been obtained from different samples of ambient noise, perhaps on different dates or from the codas of different events. The correlations are of course never identical; they are often very different. One reason for a difference is that the source of the noise may be different (yet if the correlation has converged to the local Green’s function, a change of noise source ought to have little effect). Continuous seismic sources can move and strengthen or weaken as weather changes at sea. It may also be that the correlation has not fully converged (i.e. insufficient averaging has been done). A third possibility is that the local mechanical or acoustic environment may have evolved, in particular, the local wave speed(s) may have changed. It is this possibility that is of particular interest, as changes in seismic velocities are associated with relaxations after major seismic events (Brenguier et al. 2008b). In some cases changes in seismic velocity can be used to predict volcanic eruptions (Brenguier et al. 2008a). Therefore, it is of great interest to be able to discern whether a change is due to a change in local environment or to a change in the character of the noise. The latter possibility is of some interest; the former is of great interest.

Our purpose here is to evaluate the precision with which wave speed changes can be evaluated. To do this, we consider the case in which the two waveforms \( \phi_1(t) \) and \( \phi_2(t) \) differ only by noise so that the actual relative dilation without noise, is zero. We then ask for the apparent (non-zero in general) value of \( \epsilon \) at which the corresponding \( X \) in eq. (1) achieves its maximum. The next section calculates the root mean square \( \langle \epsilon^2 \rangle \) of this apparent, and erroneous, relative dilation. The subsequent sections compare this prediction with experiment.

2 DILATION CORRELATION COEFFICIENT

Here, we examine the apparent waveform-dilation between two nominally identical signals. Theoretically, one ought to infer a relative dilation \( \epsilon \) of zero, however, noise can corrupt the inference. Key to the following analysis is an understanding that the signals being discussed are like coda, in that they are statistically stationary with durations long compared to an inverse bandwidth. We take the two waveforms to have an identical part \( \psi(t) \), and to differ by noise \( 2 \mu \chi(t) \). In the limit \( \mu \to 0 \), the waveforms become identical and have no relative dilation. If \( \mu \neq 0 \), there will be an apparent, but actually meaningless, temporal dilation between them. We wish to estimate this erroneous apparent relative dilation, and to identify any signatures that could be used to alert to the possibility of error. Note that the common part \( \psi \) of the signals need not be the local Green’s functions.

We split the difference between these two waveforms \( \phi_1 \) and \( \phi_2 \), and define two signals \( \psi \) and \( \chi \):

\[
\phi_{1,2} = \psi(t) \pm \mu \chi(t).
\]

The waveform dilation-correlation coefficient (11) between them is

\[
X(\epsilon, \mu) = \frac{\int \phi_1(t(1 + \epsilon/2)) \phi_2(t(1 - \epsilon/2)) dt}{\sqrt{\int \phi_1^2(t(1 + \epsilon/2)) dt \int \phi_2^2(t(1 - \epsilon/2)) dt}}
\]

\[
= \sqrt{1 - \epsilon^2/4} \frac{\int [\psi(t(1 + \epsilon/2)) + \mu \chi(t(1 + \epsilon/2))] [\psi(t(1 - \epsilon/2)) - \mu \chi(t(1 - \epsilon/2))] dt}{\sqrt{\left( \int \psi^2 + \mu^2 \chi^2 \right) dt} - 4 \mu^2 \left( \int \chi \psi \right)^2}.
\]

\[
= \sqrt{1 - \epsilon^2/4} \frac{N(\epsilon, \mu)}{D(\mu)}
\]

with \( N \) and \( D \) defined as, respectively, the numerator and the denominator of \( X \). The integrations are typically taken over a finite time-window with tapered edges. We make the approximation that the change of variable \( t(1 + \epsilon/2) \to t \) and \( t(1 - \epsilon/2) \to t \) in the denominator only leaves a prefactor \( \sqrt{1 - \epsilon^2/4} \).

The value of \( \epsilon \) at which \( X \) achieves its maximum is the practitioner’s estimate of the dilation between the waveforms \( \phi_1 \) and \( \phi_2 \). It occurs at \( \epsilon \) such that \( \partial X/\partial \epsilon = 0 \), or,

\[
0 = \sqrt{1 - \epsilon^2/4} D(\mu) \frac{\partial X(\epsilon, \mu)}{\partial \epsilon} = -\frac{\epsilon}{4} N(\epsilon, \mu) + (1 - \epsilon^2/4) \frac{\partial N(\epsilon, \mu)}{\partial \epsilon}.
\]

© 2011 The Authors, GJI, 185, 1384–1392
Geophysical Journal International © 2011 RAS
If the phase shift due to time dilation is much less than one oscillation, which implies \( t \omega \ll 1 \) for all times \( t \) and frequencies \( \omega \) of interest, it suffices to expand \( N(\epsilon, \mu) \) through only the second power of \( \epsilon \)

\[
N(\epsilon, \mu) = \int \left[ \dot{\psi}(t) + \frac{t \epsilon}{2} \psi(t) + \frac{t^2 \epsilon^2}{8} \ddot{\psi}(t) + \mu \chi(t) + \frac{\mu \epsilon}{2} \dot{\chi}(t) + \frac{\mu t \epsilon^2}{8} \ddot{\chi}(t) \right] \, dt.
\]

On collecting terms in \( N(\epsilon, \mu) \) that are linear and quadratic in \( \epsilon \), one obtains

\[
N(\epsilon, \mu) \sim \int \left[ \dot{\psi}(t)^2 - \mu^2 \chi(t)^2 \right] \, dt + \int \left[ \frac{t \epsilon}{2} \dot{\psi}(t) + \frac{\mu \epsilon}{2} \dot{\chi}(t) \right] \left[ \psi(t) - \mu \chi(t) \right] \, dt
\]

\[
+ \int \left[ \frac{t^2 \epsilon^2}{8} \ddot{\psi}(t) + \frac{t^2 \mu \epsilon^2}{8} \ddot{\chi}(t) \right] \left[ \psi(t) + \mu \chi(t) \right] \, dt
\]

\[
+ \int \left[ \frac{t^2 \epsilon^2}{8} \ddot{\psi}(t) - \frac{t^2 \mu \epsilon^2}{8} \ddot{\chi}(t) \right] \left[ \psi(t) - \mu \chi(t) \right] \, dt
\]

\[
+ \int \left[ \frac{t \epsilon}{2} \dot{\psi}(t) + \frac{\mu \epsilon}{2} \dot{\chi}(t) \right] \left[ - \frac{t \epsilon}{2} \dot{\psi}(t) + \frac{\mu \epsilon}{2} \dot{\chi}(t) \right] \, dt.
\]

The first term in \( \epsilon^2 \) may be integrated by parts.

\[
N(\epsilon, \mu) \sim \int \left[ \dot{\psi}(t)^2 - \mu^2 \chi(t)^2 \right] \, dt + \epsilon \int t \mu \left[ \dot{\chi}(t) \psi(t) - \chi(t) \dot{\psi}(t) \right] \, dt
\]

\[
- \frac{1}{2} \epsilon^2 \int t^2 \left[ \dot{\psi}(t)^2 - \mu^2 \dot{\chi}(t)^2 \right] \, dt + \frac{1}{4} \epsilon^2 r
\]

where the quantity \( r \) is

\[
r = \int \left[ \dot{\psi}(t)^2 - \mu^2 \dot{\chi}(t)^2 \right] \, dt - t \left[ \dot{\psi}(t)^2 - \mu^2 \dot{\chi}(t)^2 \right] - t^2 \left[ \psi(t) \dot{\psi}(t) - \mu \chi(t) \dot{\chi}(t) \right].
\]

whose expectation is zero and whose typical value is much less (by a factor of \( t^2 \omega^2 \)) than the other coefficient of \( \epsilon^2 \) in eq. (8). For this reason we henceforth neglect it.

So finally, neglecting high order of \( \epsilon \),

\[
\frac{\partial N(\epsilon, \mu)}{\partial \epsilon} \sim \int \mu t \left[ \dot{\chi}(t) \dot{\psi}(t) - \chi(t) \ddot{\psi}(t) \right] \, dt - \epsilon \int t^2 \left[ \dot{\psi}(t)^2 - \mu^2 \dot{\chi}(t)^2 \right] \, dt.
\]

Eq. (4) is satisfied for

\[
\epsilon = n/d,
\]

with

\[
n = \mu \int t \left[ \dot{\chi}(t) \dot{\psi}(t) - \chi(t) \ddot{\psi}(t) \right] \, dt,
\]

\[
d = \int \left[ t^2 \dot{\psi}(t)^2 - \mu^2 \dot{\chi}(t)^2 \right] \, dt + \frac{1}{4} \int \left[ \dot{\psi}(t)^2 - \mu^2 \dot{\chi}(t)^2 \right] \, dt.
\]

Eq. (11) is an expression for the apparent dilation induced by the difference \( 2 \mu \chi \) between the original waveforms \( \phi_1 \) and \( \phi_2 \). Given specific \( \psi \) and \( \chi \), one could evaluate it. It will be more useful, however, to obtain statistical estimates for the apparent dilation given assumptions about the envelopes and spectra of \( \psi \) and \( \chi \). The numerator \( n \) has expectation zero, as \( \chi \) and \( \psi \) are statistically unrelated. Thus within the stated limit \( \omega \epsilon \ll \omega \), differences \( \phi_2 - \phi_1 \) do not manifest as an apparent dilation and the expected dilation \( \epsilon \) is zero.

### 3 Variance Estimation and Statistical Error

Given \( n = 0 \), one then seeks estimates for the root mean square of eq. (11) to judge typical fluctuations around the expected zero. These will be made based on assumptions that \( \psi \) and \( \chi \) are stationary, noise-like and Gaussian, with similar spectra, having central frequency \( \omega_0 \).
\(\psi\) and \(\chi\) have the same duration, long compared to the inverse of \(\omega_c\). Without loss of generality it is also assumed that they have the same amplitudes \((\psi^2) = (\chi^2) = 1\). They are taken to extend from a start time \(t_1\) to an end time \(t_2\). Under these assumptions the denominator of (11) is estimated as

\[
d \approx (1 - \mu^2) \left[ \frac{1}{3} \omega_c^2 (t_2^3 - t_1^3) + \frac{1}{4} (t_2 - t_1) \right] \approx (1 - \mu^2) \frac{1}{3} \omega_c^2 (t_2^3 - t_1^3).
\]

The square of the numerator of (11) is

\[
n^2 \approx \mu^2 \left[ \int \int \tau \{ \psi(t) \dot{\chi}(t) - \dot{\psi}(t) \chi(t) \} \{ \dot{\psi}(t') \chi(t') - \dot{\psi}(t') \chi(t') \} \, dt \, dt' \right].
\]

On changing variables: \(t + t' = 2\tau, t - t' = \xi\) and dropping the cross terms as having expectation zero, (13) becomes

\[
\langle n^2 \rangle \approx \mu^2 \int \left( \tau^2 - \frac{\xi^2}{4} \right) \{ \psi(t + \xi) \dot{\chi}(t) \} \{ \dot{\psi}(t) \chi(t) \} \, dt \, d\xi.
\]

Autocorrelation functions may be defined as

\[
\langle \psi(t + \xi) \dot{\psi}(t - \xi) \rangle = \langle \psi^2(\tau)R_\psi(\xi) \rangle = R(\xi),
\]

such that

\[
\langle \psi(t + \xi) \chi(t - \xi) \rangle \approx \omega_c^2 \langle \dot{\psi}^2(\tau) \rangle R_\psi(\xi) = \omega_c^2 R(\xi),
\]

with similar expressions for \(\chi\). Then the expectation of the square of the numerator of (11) is

\[
\langle n^2 \rangle \approx 2 \mu^2 \left[ \int \left( \tau^2 - \frac{\xi^2}{4} \right) \omega_c^2 R^2(\xi) \, dt \, d\xi \right] \approx 2 \mu^2 \omega_c^2 \left[ \int \tau^2 \, dt \right] \left[ \int R^2(\xi) \, d\xi \right].
\]

The first integral is merely \((t_2^3 - t_1^3)/3\). The second requires knowing some of the spectra of \(\psi\) and \(\chi\), so we take these to be Gaussian and identical: \(\sim \exp(-\omega^2/2T^2)\). \(T\) may be identified by noting that the \(-10\) dB points are at \(\omega_c\) plus or minus in \(10T\). In this case, \(R\) is related to the inverse Fourier transform of the power spectrum \(R(\xi) = \cos(\omega_c \xi) \exp(\xi^2/4T^2)\). Then the second integral in (17) is identified as \(T/\sqrt{\pi}T\).

Application of eqs (11), (12), (17) requires that we also estimate the \(\mu\). The quantity \(\mu\) is related to the maximum of the waveform dilation-correlation coefficient

\[
X(0, \mu) = \frac{N(0, \mu)}{D(\mu)} = \frac{\int \psi(t)^2 - \mu^2 \chi(t)^2 \, dt}{\sqrt{\int \psi^2 + \mu^2 \chi^2 \, dt^2 - 4\mu^2 \int \chi \psi \, dt^2}}.
\]

As \(\chi\) and \(\psi\) are statistically independent, one estimates the following relation between the maximum of the dilation correlation coefficient and the parameter \(\mu\)

\[
X = \frac{1 - \mu^2}{1 + \mu^2}.
\]

Finally, the root mean square of the practitioner’s (erroneous) estimate for the relative dilation between \(\phi_1\) and \(\phi_2\) is

\[
\text{rms} \, \epsilon = \frac{\langle n^2 \rangle^{1/2}}{d} = \frac{\sqrt{1 - X^2}}{2X} \sqrt{\frac{6\sqrt{\pi}T}{\omega_c^2 (t_2^3 - t_1^3)}}.
\]

We recall that \(T\) is the inverse of the frequency bandwidth, \(t_1\) and \(t_2\) are begin and end time of the processed time-window in the coda, respectively and \(\omega_c\) is the central pulsation. This expression scales inversely with the duration of the correlation waveform in units of the period, and inversely with the square root of the duration in units of the inverse bandwidth. In practice eq. (20) can be very small. The quantity \(\omega_c\) represents the available time where coda waves are significantly larger than the noise; the duration of the waveform is in units of the period. The quantity \(T\) is the amount of time for one bit of information to be delivered, and corresponds roughly to the time of the initial source (Derode et al. 1999). Thus eq. (20) can be recognized as scaling inversely the available time in the coda (this time is related to coda-Q), and inversely with the square root of the amount of information. It also may be recognized that small \(X\) corresponding to waveforms \(\phi_1\) and \(\phi_2\) that are very different, permits the practitioners erroneous estimate of dilation to be large. It may be that lengthening the considered time interval \(t_2 - t_1\) would increase the precision, however it could also diminish \(X\): in principle there are trade-offs.

Application of eq. (20) is straightforward. A practitioner’s estimate of the relative dilation \(\epsilon\) between two waveforms \(\phi_1\) and \(\phi_2\) may be compared to eq. (20). Values in excess of eq. (20) are consistent with the inference that the observed dilation is real. Changes in waveform source or other character should not generate apparent dilations in excess of eq. (20). Furthermore, in absence of any actual dilation, estimates of \(\epsilon\) of the order eq. (20) will nevertheless be generated in practice. Such should be regarded as unmeaningful.
4 Comparison with Experiment

The prediction eq. (20) has been compared to waveform dilation measurements in a laboratory ultrasonic experiment (Hadziioannou et al. 2009). For practical reasons, we mimicked here ambient noise correlation with diffuse waves correlation. Several piezoelectric sensors and sources were applied to a multiple scattering air-bubble filled gel (see Fig. 1). Sources and receivers were placed on opposite sides, 64 mm apart. Multiple scattering was strong: received waveforms \( f_{sr}(t) \) from sources \( s \) to receivers \( r \) were coda-like, with envelopes that resembled the solution of a diffusion equation (Fig. 2). The autocorrelation of each \( f_{sr}(t) \) was windowed between lapse times of 12.5–50 \( \mu s \), to yield the waveforms which we call \( g_{sr}(\tau) \) (see Fig. 3). Details of the experimental set-up are described in Hadziioannou et al. (2009). The details are, however, unimportant here, as the present theory applies to any pair of coda-like waveforms \( \phi_1 \) and \( \phi_2 \). The typical \( g_{sr} \) is stationary over this interval and has a power spectrum centred on 2.35 MHz with \(-10\) dB points at 1.7–3.0 MHz.

The tables below are formed by maximizing the dilation correlation coefficient \( X \) between sums \( \phi = \sum_g g_{sr} \) over different sets of sources \( (s) \). Note that the \( \phi \) are not Green’s functions \( G_{rr} \), as the fields \( f_{sr}(t) \) used to compose them were not fully equipartitioned. The excellent impedance match between the gel and the receivers prevented the field to be reflected back to the medium. The noise field thus lacked any components travelling from receiver to receiver. All tests were conducted at fixed temperature, the actual relative temporal dilation is therefore zero. Also, to mimic signals acquired at different dates, we averaged correlations over different set of sources to eventually compare them. The addition of different sources results in an additional noise \( \chi(t) \) in the correlations. The goal of the test is to measure the dilation induced by the difference in waveform due to a different source distribution.

Autocorrelation waveforms, like that illustrated in Fig. 3, in the interval from 12.5 to 50 \( \mu s \) appear stationary. Thus we take \( t_1 = 12.5 \mu s \), \( t_2 = 50 \mu s \), \( \omega_c = 15 \text{ rad}\mu s^{-1} \) and \( T = 0.56 \mu s \) and conclude from (20),

\[
\text{rms } \epsilon = 4 \times 10^{-4} \sqrt{1 - X^2}.
\]

(21)

Tables 1 and 2 show two case studies. In the first case, autocorrelations calculated from the signals at a receiver \( r \), as produced by 11 distinct sources \( s \), were summed over to generate the reference waveform \( \phi_1 = \sum_s g_{sr} \). For each of three comparison waveforms \( \phi_2 \), the same sum was done, keeping the first 10 sources unchanged. To deliberately change the waveform without dilation, the eleventh source is replaced

![Figure 1](https://example.com/figure1.png)

**Figure 1.** Experimental setup with ultrasound. An air-gel mixture mimics a multiple scattering medium. Coda waves sensed by the receivers (see Fig. 2) are processed like ambient seismic noise: they are autocorrelated and compared from one date to another.

![Figure 2](https://example.com/figure2.png)

**Figure 2.** A typical signal \( f_{sr}(t) \) in the gel resembles noise, under an envelope which is a solution of a diffusion equation.
Figure 3. A typical autocorrelation $g_{ss}$ of the signal from one of the sources to one of the receivers. The interval from 12.5 to 50 $\mu$s was selected for dilation coefficient evaluation.

Table 1. Comparison of best-fit waveform dilations $\epsilon$ with the predictions of eq. (21). A maximum value of $X$ and the $\epsilon$ at which that $X$ is maximum, are constructed for each of seven receivers (the seven columns) and the three choices for the set of sources described in the text (the three rows). The root mean square of those $\epsilon$ is compared with the predictions of theory. That $X$ is of order 90 per cent is consistent with one source in ten having changed.

<table>
<thead>
<tr>
<th>Sources 1–11 and 12</th>
<th>0.9312</th>
<th>0.9169</th>
<th>0.8893</th>
<th>0.8458</th>
<th>0.8226</th>
<th>0.8852</th>
<th>0.8652</th>
<th>0.8683</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sources 1–11 and 13</td>
<td>0.9394</td>
<td>0.8833</td>
<td>0.9083</td>
<td>0.8464</td>
<td>0.8872</td>
<td>0.8631</td>
<td>0.8928</td>
<td></td>
</tr>
<tr>
<td>Sources 1–11 and 14</td>
<td>0.9458</td>
<td>0.9009</td>
<td>0.8730</td>
<td>0.7942</td>
<td>0.8322</td>
<td>0.8396</td>
<td>0.7979</td>
<td></td>
</tr>
</tbody>
</table>

The dilation $\epsilon$ ($\times 10^{-3}$) as obtained by maximizing $X$ for each of these cases

<table>
<thead>
<tr>
<th>Sources 1–11 and 12</th>
<th>0.06</th>
<th>0.04</th>
<th>-0.16</th>
<th>0.08</th>
<th>-0.04</th>
<th>-0.10</th>
<th>0.18</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sources 1–11 and 13</td>
<td>-0.04</td>
<td>-0.04</td>
<td>-0.08</td>
<td>0.06</td>
<td>-0.14</td>
<td>-0.08</td>
<td>0.10</td>
</tr>
<tr>
<td>Sources 1–11 and 14</td>
<td>-0.16</td>
<td>0.08</td>
<td>-0.12</td>
<td>0.00</td>
<td>-0.14</td>
<td>-0.24</td>
<td>-0.12</td>
</tr>
</tbody>
</table>

Experimental root mean square dilation $\epsilon$ ($\times 10^{-3}$)

<table>
<thead>
<tr>
<th>All sets</th>
<th>0.1013</th>
<th>0.0566</th>
<th>0.1244</th>
<th>0.0577</th>
<th>0.1166</th>
<th>0.1571</th>
<th>0.1376</th>
</tr>
</thead>
</table>

Theoretical root mean square ($\times 10^{-3}$) from eq. (21)

<table>
<thead>
<tr>
<th>All sets</th>
<th>0.07</th>
<th>0.09</th>
<th>0.09</th>
<th>0.12</th>
<th>0.11</th>
<th>0.10</th>
<th>0.11</th>
</tr>
</thead>
</table>

Table 2. As in Table 1, but for sources that differ by more: fewer sources are kept fixed (four) and more sources are changing (two). This results in a smaller values of $X$ and a larger value of error (uncertainty).

<table>
<thead>
<tr>
<th>Sources 1–4 and 5,6</th>
<th>0.6181</th>
<th>0.7864</th>
<th>0.7143</th>
<th>0.8400</th>
<th>0.7149</th>
<th>0.8458</th>
<th>0.7863</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sources 1–4 and 7,8</td>
<td>0.6359</td>
<td>0.7340</td>
<td>0.7011</td>
<td>0.8451</td>
<td>0.8020</td>
<td>0.8285</td>
<td>0.8194</td>
</tr>
<tr>
<td>Sources 1–4 and 9,10</td>
<td>0.5948</td>
<td>0.7397</td>
<td>0.5837</td>
<td>0.8165</td>
<td>0.8294</td>
<td>0.8451</td>
<td>0.8745</td>
</tr>
</tbody>
</table>

The dilation $\epsilon$ ($\times 10^{-3}$) as obtained by maximizing $X$ for each of these cases

<table>
<thead>
<tr>
<th>Sources 1–4 and 5,6</th>
<th>-0.0800</th>
<th>0.0400</th>
<th>0.4600</th>
<th>-0.0200</th>
<th>0.1400</th>
<th>0.0400</th>
<th>0.1600</th>
</tr>
</thead>
<tbody>
<tr>
<td>Sources 1–4 and 7,8</td>
<td>-0.0200</td>
<td>0.0400</td>
<td>0.0400</td>
<td>-0.3400</td>
<td>-0.0800</td>
<td>-0.1400</td>
<td>0.0800</td>
</tr>
<tr>
<td>Sources 1–4 and 9,10</td>
<td>-0.4600</td>
<td>0.1600</td>
<td>0.5600</td>
<td>-0.0400</td>
<td>0.0800</td>
<td>-0.0400</td>
<td>0.0400</td>
</tr>
</tbody>
</table>

Experimental root mean square dilation $\epsilon$ ($\times 10^{-3}$)

<table>
<thead>
<tr>
<th>All sets</th>
<th>0.27</th>
<th>0.098</th>
<th>0.419</th>
<th>0.198</th>
<th>0.104</th>
<th>0.087</th>
<th>0.106</th>
</tr>
</thead>
</table>

Theoretical root mean square ($\times 10^{-3}$) from eq. (21)

| All sets | 0.19 | 0.14 | 0.17 | 0.12 | 0.14 | 0.12 | 0.11 |
with sources number 12, 13 and 14, respectively. This was repeated for each of seven receivers. In each case, we compare three waveforms \( \phi_2 \) with the reference \( \phi_1 \) and evaluate \( X(\epsilon) \). The table shows the maximum value of \( X(\epsilon) \), and the value of \( \epsilon \) that did this, for each of the 21 cases. For each of the seven receivers we calculate the rms of these three \( \epsilon \). If the only changes were to the source of the noise field, and not the medium, one would expect no dilation, or \( \epsilon = 0 \). Nevertheless, the differences in sources do generate apparent (feeble but notable) dilations \( \sim \epsilon \). Theoretical and experimental rms(\( \epsilon \)) are of the same order of magnitude. Theory, especially in light of the approximate modelling of the spectrum, may be said to have done a good job predicting the fluctuations.

In the second study (Table 2), four sources were held constant, and two were varied. Here, the reference waveform was constructed from a sum over six sources \( \sum g_{sr} \); each of the other three waveforms was constructed by replacing sources number five and six in that sum with two others. Again theory may be said to have done a good job: the rms theoretical predictions accurately fit the actual experimental errors within 40 per cent. This means that eq. (20) properly predicts the order of magnitude of the error.

5 COMPARISON WITH SEISMIC DATA FROM PARKFIELD

We also analyse data from seismic measurements near Parkfield, California. Brenguier et al. (2008b) showed that correlation waveforms obtained from ambient seismic noise over a period of 5 yr from 2002 to 2007 changed in a manner consistent with a decrease of the seismic velocity after the earthquake of 2004 (Fig. 4). This decreased velocity then relaxed like \( \log(t) \) after the earthquake. While they used the doublet technique, we have reanalysed their data using the dilation coefficient (see eq. 11). For each of 78 receiver pairs, we compared the 1550 d average correlation waveform with the correlation waveforms constructed from each of 1546 overlapping 5 d segments. The whitening operation before correlation ensures that the spectrum of the correlations is constant. Note that direct arrivals are never processed. A representative correlation waveform is shown in Fig. 5. Each such waveform was windowed between \(-50\) and \(-20\) s, and again from 20 to 50 s (thus excluding direct Rayleigh arrivals and emphasizing the multiple scattered diffuse part of the signal for which the theory was developed). As in the previous section, the details of the measurements are available elsewhere (Brenguier et al. 2008b) but are unimportant for the present purposes. An \( X \) and an \( \epsilon \) were deduced for each day. Power spectra were centred on 0.5 Hz, with \(-10\) dB shoulders at 0.1–0.9 Hz. These numbers permit the evaluation of (20)

\[
\text{rms } \epsilon = 2.4 \times 10^{-3} \frac{\sqrt{1 - X^2}}{2X} \tag{22}
\]

Fig 6 (top and bottom) shows the mean (over the 78 receiver pairs) values of \( X \) and \( \epsilon \) between each of the 1546 overlapping 5 d correlation waveforms, \( \phi_1 \) and the correlation waveform \( \phi_2 \) as obtained by averaging over the entire 5 yr period. Except for the two events on days 152 and 437, and the slow relaxation after the latter, the dilation appears constant, with daily random fluctuations of order \( 10^{-4} \). A correlation coefficient \( X \) of 0.8 predicts a rms fluctuation of \( 10^{-3} \) (eq. 22). On averaging over 78 pairs, this prediction is reduced by a factor \( \sqrt{78} \), to

![Figure 4](https://academic.oup.com/gji/article-abstract/185/3/1384/604014/120)
Precision of noise correlation interferometry

Figure 5. A typical daily correlation waveform from the Parkfield data set. Dilations were constructed by comparing waveforms like this as windowed from 50 to +50 s, with the direct signal from 20 to +20 omitted.

Figure 6. Top-panel: Dilation \( \epsilon \) averaged over 78 receiver pairs, using a 5 d sliding window. The grey squares indicate the 70 d reference windows. The best fit \( \epsilon \) varied weakly and stochastically over this period, with two notable jumps, after 2003 December 22 and after the Parkfield earthquake on 2004 September 28. The latter jump was followed by a slow recovery. Fluctuations have an rms strength of about \( 10^{-4} \). Bottom-panel: Dilation coefficient \( X \), the maximum value of the dilation coefficient \( X \), averaged over 78 receiver pairs.

1.1 \( \times 10^{-4} \), consistent with the observed fluctuations in \( \epsilon \). In light of the approximations, in particular that of modelling the spectrum as Gaussian and the waveform as stationary, we count this as excellent agreement.

The discontinuities in \( \epsilon \) at 2003 December 22 and 2004 September 28 are of particular interest. The latter is coincident with the Parkfield earthquake. Jumps in dilation on those dates by \( \sim 0.8 \times 10^{-3} \) were interpreted (Brenguier et al. 2008b) as decrease of local seismic wave speed. But one might wish to entertain the hypothesis that these jumps are due to a change in the source of the noise. To examine the question, we evaluated \( X \) and \( \epsilon \) using correlation waveforms \( \phi_1 \) as averaged over a 70 d period before each event as a reference and correlation waveforms \( \phi_2 \) obtained over a series of 5 d spans after the events. The relative dilation across the events are the same as seen in Fig. 6 (top), of order \( 5 \times 10^{-4} \). The values of \( X \) for these pairs of waveforms varied between 60 per cent and 70 per cent. According to eq. (22) divided by \( \sqrt{78} \), the value of \( X \) would have had to be below 33 per cent if this large and apparent dilation were to be due to a random function with no actual dilation. The relative dilation between correlation waveforms before and after the event is therefore due to changes in seismic Green’s function, and not to changes in the source of the waves.

6 SUMMARY

Waveforms constructed by noise correlation can be extraordinarily sensitive to changes in material properties. Such waveforms are in principle affected by both changes in noise sources and changes in the acoustic properties of the medium in which the waves propagate. It has been shown here that long-duration diffuse waveforms permit changes in the source of the noise to be distinguished with high precision from changes due to a temporal dilation.
An expression was derived for the rms of the apparent dilation $\epsilon$ measured on two waveforms, when there is no actual dilation between the two. This apparent dilation can be an effect of, for example, a change in noise sources. The rms value thus allows us to distinguish between an erroneous dilation measurement due to waveform change, and a physical wave speed change in the medium.

We have tested the validity of the rms value using data from laboratory experiments, and we find that the theory predicts errors well.

**ACKNOWLEDGMENTS**

Most of this work was done while the first author was a visitor at the Laboratoire de Géophysique Interne et Tectonophysique (Université J. Fourier). This work was partially funded by the ANR JCJC08 SISDIF grant, and by the ERC Whisper Advanced grant. All the seismic data used in this study came from the Parkfield HRSN and were collected by the Berkeley Seismological Laboratory (BSL).

**REFERENCES**


