Recycling of archaeal biomass as a new strategy for extreme life in Dead Sea deep sediments

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ABSTRACT
Archaea and Bacteria that inhabit the deep subsurface (known as the deep biosphere) play a prevalent role in the recycling of sedimentary organic carbon. In such environments, processes that occur over millions of years and require microbial communities to cope with extremely limited sources of energy. Because of this scarcity, metabolic processes come at a high energetic cost, but the ways heterotrophic microbial communities develop to minimize energy expenses for a maximized yield remain unclear. Here, we report molecular biomarker evidence for the recycling of archaeal cell wall constituents in extreme evaporitic facies of Dead Sea deep sediments. Wax esters derived from the recombination of hydrolyzed products of archaeal membrane lipids were observed in gypsum and/or halite sedimentary deposits down to 243 m below the lake floor, implying the reutilization of archaeal necromass possibly by deep subsurface bacteria. By recycling the building blocks of putatively better-adapted archaea, heterotrophic bacteria may build up intracellular carbon stocks and mitigate osmotic stress in this energy-deprived environment. This mechanism illustrates a new pathway of carbon transformation in the subsurface and demonstrates how life can be maintained in extreme environments characterized by long-term isolation and minimal energetic resources.

INTRODUCTION
In extreme environments, microbial metabolic processes that lower the energetic cost of maintaining life are favored (Hoehler and Jørgensen, 2013). Such settings are characterized by low growth rates (Lomstein et al., 2012), and most energy is diverted to maintenance functions (van Bodegom, 2007), such as osmotic equilibration, oxygen stress defense, motility, and more sustainable metabolic pathways. These selective conditions for life promote the dominance of prokaryotes and generally favor Archaea relative to Bacteria (Hoehler and Jørgensen, 2013). This is mostly due to the reduced membrane permeability of Archaea, which requires less maintenance energy with respect to bacterial membranes (Valentine, 2007). This advantage is particularly striking in environments characterized by high osmotic stress such as hypersaline environments. There, bacteria may use alternative strategies that allow competition with the putatively better-adapted archaea, for example, by recycling available organic molecules as osmotic solutes (Oren, 1999a). The intracellular accumulation of ambient organic carbon is a common way of economizing energy in harsh environments. For example, under stressed growth conditions, some Bacteria species are known to accumulate intracellular lipid droplets (Alvarez et al., 1997; Wältermann and Steinbüchel, 2005) in the form of polyhydroxyalkanoates, triglycerides, or wax esters (WEs). However, the presence of such mechanisms in the deep biosphere has not yet been documented, suggesting that they may not be sufficiently effective in these low-energy settings for bacteria to survive.

The Dead Sea is the most saline lake on Earth and has deposited evaporitic minerals since the early Quaternary (Stein, 2001). As a result, its subsurface environment constitutes one of the most extreme ecosystems on the planet. The extreme chemistry of the water allows only for the survival of halophilic archaea in the water column and recent halite sediments (Bodaker et al., 2010; Thomas et al., 2015). Bacteria have rarely been observed in the most extreme sedimentary facies of the Dead Sea (halite or gypsum), suggesting that these harsh conditions limit their growth (Thomas et al., 2015). To further investigate the composition of the microbial community and its potential metabolic strategies to survive the most arid periods of the late Quaternary of the Dead Sea basin, we characterized the lipid biomarker composition of deep hypersaline halite and gypsum sediments. The good lipid preservation in these extreme horizons gives clues to metabolic pathways that allow for the survival of a thus-far-unrecognized deep biosphere.

METHODS
The material used in this study originated from the International Continental Scientific Drilling Program (ICDP) Dead Sea Deep Drilling Project (DSDDP) core 5017–1A retrieved from the center of the Dead Sea in winter 2010–2011 (Fig. 1). Sediments were sampled on site from core catchers using sterile tools and kept in the freezer at –20 °C until further processing. The main characteristics of samples analyzed for lipid biomarkers are given in Table DR1 in the GSA Data Repository1. Samples were freeze-dried, ground, and extracted using multiple sonication cycles (methanol 2×, methanol/dichloromethane [DCM] [1:1 v/v], methanol/dichloromethane [DCM] [1:1 v/v], DCM 3×). Elemental sulfur was removed with activated copper. Lipid extracts were filtered and separated using a deactivated silica gel column (5% H2O) into five fractions of increasing polarity (hexane/DCM [9:1 v/v], hexane/DCM [1:1 v/v], DCM, ethyl acetate, and methanol). Fractions 3 and 4 were silylated with pyridine/bis(trimethylsilyl) trifluoroacetamide (BSTFA) at 2:1 (v/v). Fraction 5 was trans-esterified by overnight incubation with 0.5 mL of toluene and 2 mL of 2% Na2SO4 in methanol at 60 °C. Following the addition of NaCl (5%), the organic phase was extracted with hexane:DCM (4:1 v/v, 3×), washed with NaHCO3 (2%), and dried with sodium sulfate before silylation.

All fractions were analyzed by gas chromatography mass spectrometry (GC-MS) on an HP 6890 Series Plus gas chromatograph equipped

1GSA Data Repository item 2019179, supplementary Figures DR1 and DR2 and Tables DR1 and DR2, is available online at http://www.geosociety.org/datarepository/2019/, or on request from editing@geosociety.org.

with a cool on-column injector, and coupled to an Agilent 5975C (VL MSD) mass spectrometer. The GC was equipped with an HP5 column (30 m × 0.25 mm × 0.25 µm, RESTEK). Samples were injected at 60 °C (held for 30 s), and the oven temperature was increased to 130 °C at a rate of 20 °C/min, then to 250 °C (5 °C/min) and 300 °C (3 °C/min), and finally held isothermal for 45 min. Stepwise dilution of external standards allowed quantification. Isoprenoid fatty acids are less polar than linear acids and partly eluted in fraction F4. A ratio based on specific bacterial fatty acids was calculated using the sum of methyl-branched C_{15} and C_{17} fatty acids over the sum of linear C_{15} and C_{17} fatty acids. Changes in this ratio indicate a shift in bacterial community.

Compound-specific carbon isotope (δ^{13}C) analyses were done using a HP7890B GC coupled to an isoprime visION isotope ratio mass spectrometer via a GC-5 combustion interface coupled to an Isoprime visION isotope ratio mass spectrometer (Arndt Schimmelmann, Indiana University, and the oven temperature was programmed as for GC-MS analyses. The B4 standard mixture (Arndt Schimmelmann, Indiana University, USA) was used to externally calibrate the δ^{13}C values, and the known BSTFA derivatizing agent δ^{13}C value was corrected for alcohols and fatty acids. After samples were decaffeinated, total organic carbon (TOC) was measured, and data were normalized to a standard sediment (IVA Analysezentrum, Germany) containing 9.15% of carbon, measured every 10 samples.

RESULTS AND DISCUSSION

We found significant amounts of isoprenoid wax esters (WEs; up to 0.2 µg g^{-1} TOC) in halite and gypsum samples retrieved between 90 and 250 m below lake floor (mbfl; Fig. 2; Table DR1). WEs are an important type of energy-storage molecule that can be produced by Eukarya and Bacteria, particularly under conditions of stress, but they have not been reported in the archaeal domain so far (Wang et al., 2019). WEs are formed by condensation of fatty acids and alcohols available in the environment and provide easily accessible (intracellular) sources of carbon (Waltermann and Steinbüchel, 2005). The major WE detected was phytanyl phytanate (3,7,11,15-tetramethylhexadecyl-3,7,11,15-tetramethylhexadecanooate; iC_{20}-iC_{20}; Fig. 3), which was accompanied by a series of other isoprenoid WEs composed of a C_{15} or C_{17} isoprenoid alcohol esterified to a linear, methyl-branched or isoprenoid acyl chain (Fig. 3A). The isoprenoid WEs were accompanied by significant amounts of membrane core lipids of halophilic archaea—archaeol (2,3-O-diphytanyl-1-sn-glycerol) and extended archaeol (2,3-O-phytanyl-1-O-sesterterpapany1-sn-glycerol; Kates, 1997; Dawson et al., 2012)—along with several of their hydrolyzed and eventually oxidized products (Figs. 2 and 3B; Table DR2). Lipids of halophilic archaea are composed of C_{30} and C_{35} isoprenoid alkyl chains, which allow for better control of membrane permeability under strong osmotic stress compared to bacterial fatty acid membranes (Valentine, 2007; Koga, 2012). Archaeol was the most abundant isoprenoid alcohol in most of the investigated sediment samples (Figs. 2 and 3). Extended archaeol, a membrane lipid more specific to halophilic archaea of the Halobacteria class (Dawson et al., 2012) that dominate the Dead Sea halite and gypsum sediments (Thomas et al., 2015), was also found to be enriched in halite and gypsum facies. In most of the sedimentary intervals where WEs occurred, archaeol and extended archaeol were found in lower abundance than in the other halite/gypsum samples (Fig. 2). All WE-containing intervals also contained hydroxylated core lipids and some oxidized counterparts (Fig. 2), which consisted of C_{30} and C_{35} isoprenoid alcohols and acids (phytanol, phytic acid, and C_{30} homologues), and of isoprenoid C_{30} and C_{35} monoalkyl glycerol monooethers (Fig. 3B; Table DR2). The report of WEs with isoprenoid C_{35} carbon chains (iC_{20}-iC_{20} and iC_{20}-iC_{20}) in environmental samples is unprecedented and attests to the incorporation of archaeal core lipid subunits into isoprenoid WEs. The formation of isoprenoid WEs has been previously demonstrated during growth of bacteria on free isoprenoid compounds such as phytane, phytol, or squalene (Rontani et al., 1999, 2003; Silva et al., 2007). The recycling of phytol derivatives into mixed isoprenoid and non-isoprenoid WEs has also been reported in desiccated mats from modern evaporative alkaline lakes (Finkelstein et al., 2010). However, phytane, phytol, and squalene were not observed in the investigated samples from the Dead Sea. As a result, the WEs detected in the Dead Sea halite/gypsum samples mostly consist of lipid subunits derived from the cell walls of archaea, the dominant organisms in the
suggestions a bacterial origin of the WEs. Recent work on fluid inclusions allowed the retrieval of bacterial sequences in halite-dominated layers of the deep Dead Sea sediment (Thomas and Ariztegui, 2019). Our data would therefore support the development of bacteria upon archaeal necromass in the extreme environment of the Dead Sea subsurface.

The concentration of intracellular WEs in bacteria has been shown to arise particularly in situations of nutrient starvation, especially when nitrogen is limiting (Ishige et al., 2003). Additionally, the recycling of archaeal biomass by archaea themselves was suggested to minimize the energy cost of life in nutrient- and carbon-limited environments (Takano et al., 2010). Here, we propose that the recycling of archaeal necromass into storage lipids like WEs constitutes a way to save and store energy in the nutrient- and energy-demanding hypersaline environment of the Dead Sea. The presence of WEs with mixed linear and isoprenoid acyl chains alongside the predominant isoprenoid WEs indicates that the WE-forming microbial population also recycled bacterial and/or eukaryal fatty acids, in addition to archaeal lipids (Fig. 4). This evidence suggests that the subsurface biosphere creates easily accessible carbon stocks from necromass, in a lake with rare allochthonous inputs and primary production (Oren, 1999b).

The accumulation of WEs has previously been reported in episodically desiccated microbial mats in hypersaline environments (Finkelmstein et al., 2010), and this was suggested to allow for better microbial cell survival during periods of desiccation, due to the release of $H_2O$ during esterification. The Dead Sea sediments that bear WEs originate from the deepest part of the lake (i.e., 297.5 m below lake level), where neither traces of desiccation (Neugebauer et al., 2016) nor microbial mats have been reported. These sedimentary levels correspond, however, to

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**Figure 3.** Characteristic chromatograms of isoprenoid wax esters (WEs) and related building blocks. A: Total ion chromatogram (TIC) of WE fraction (fraction 2) from gypsum sample at 91.04 m below lake floor (mblf). B: Extracted ion chromatogram (m/z 103+133+159+205+218) of a polar lipid fraction (fraction 4) from gypsum sample at 243.22 mblf. These intervals were selected because they show highest concentrations of isoprenoid WEs and hydrolyzed core lipids, respectively. MGM—monoalkyl glycerol monoethers. Red symbols indicate an archaeal origin; blue symbols indicate a bacterial/eukaryal origin. Numbers above symbols shown in B correspond to number of carbon atoms in chain.

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**Figure 4.** Schematic pathway for isoprenoid wax ester (WE) formation by recycling of membrane lipids in Dead Sea sediments. Intact phospholipids from archaeal and bacterial/eukaryal communities are hydrolyzed and eventually oxidized, and some of these degradation products are transformed into WEs by bacterial communities in deep halite/gypsum sediment horizons. Symbols refer to compounds identified in Figure 3. Red moieties have an archaeal origin; blue moieties have a bacterial/eukaryal origin.
periods when salinity was the highest in the deep brine, as supported by the bromide concentration curve (Fig. 2), a reliable salinity proxy for the Dead Sea water column during the Quaternary (Levy et al., 2017). The WE-bearing levels generally correspond to the highest concentrations of this element (Fig. 2; Table DR1), which are observed during the driest climatic intervals (marine isotopic stages 5E, early 5A, and the 2/1 transition). Water availability is a major issue, not only in dry environments, but also in hypersaline settings due to high salt concentrations inducing osmotic stress. By recombining hydrolyzed moieties of core lipids originating from the buried archaeal necromass, the WE-forming bacterial population may therefore have created accessible water molecules that favored its survival.

Our results illustrate the high adaptability of the subsurface biosphere and its ability to use varied strategies for energy production and preservation under adverse conditions. By studying an environment that pushes life to its limits, we catch a glimpse of the processes that fuel life in the deep subsurface, and we advance our understanding of the deep sedimentary carbon cycle.

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