

Did intense volcanism trigger the first Late Ordovician icehouse?: COMMENT

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Details of the climate evolution during the Middle and Late Ordovician are increasingly debated, with implications for understanding both greenhouse-icehouse transitions and the Late Ordovician mass extinction. We applaud Buggisch et al. (2010) for attempting to distinguish among different alternatives of climate states during this time period using temperature estimates from the relatively robust phosphate-oxygen paleothermometer as recorded in conodonts and fish teeth. However, we think they have made an error in plotting the position of their Minnesota samples relative to the Deicke K-bentonite (DKb). This error does not diminish the importance of their data, but it does remove support for their conclusion that the volcanic eruption that led to the deposition of the DKb initiated a short-term cooling event of at least 6 °C, and possibly a first short-term glacial episode during the Late Ordovician.

Buggisch et al. plotted none of their samples below the DKb (their figures 1 and 2). A critical review of the original stratigraphic description of the location of the conodont samples for the Minnesota section (Thompson, 1959; Webers, 1966), as described in their supplementary data set (Table DR1 in GSA Data Repository item 2010087), indicates that eight samples predate the Deicke eruption. This includes the sample with the highest $\delta^{18}\text{O}$ value (Fig. 1). Samples with sample letters "PI" (see Table DR1 in Buggisch et al.'s Data Repository) are from the Cummingsville Annex Section (CAS), which has been described in detail by several researchers (Sloan, 1987; Thompson, 1959; Webers, 1966). Specifically, we argue that samples PI10/11, PI12, PI13, both of the PI14 samples, PI16, and PI16A are from below the DKb, with sample PI16 (described as collected from a shale that is associated with a bentonite [Webers, 1966]) occurring directly below the DKb, because

1. The bentonite has been identified and described as the DKb (Sloan, 1987).

2. A shale horizon is a common occurrence in the upper Midwest United States below the DKb (e.g., Weiss, 1955) including the CAS, where a shale interval is present immediately underneath the DKb but not above it (Thompson, 1959).

3. Samples PI16 and below contain *Polyplacognathus*, which in Minnesota only occurs in significant numbers below the DKb (Leslie, 2000, 2009; Sloan, 1987; Webers, 1966; and our own observations). Sample PI14 (and PI14A) is rich in *Polyplacognathus* sp., and the faunas listed by Thompson (1959) in the shale horizon below the DKb match the faunas reported by Webers (1966) in terms of species (note that Thompson [1959, p. 67] thought *Polyplacognathus* was a synonym of *Amorphognathus*).

4. Samples PI10/11, PI12, and PI13 are from the McGregor Member of the Platteville Formation. The McGregor Member underlies the Carimona Member of the Platteville Formation in Minnesota (Sloan, 1987; Webers, 1966). Thus, the samples from the McGregor Member all predate the DKb.

Replotting the data with correct position relative to the DKb (Fig. 1) shows that the evidence for cooling precedes the deposition of the bentonite. Therefore, the eruption could not have triggered cooling. In other

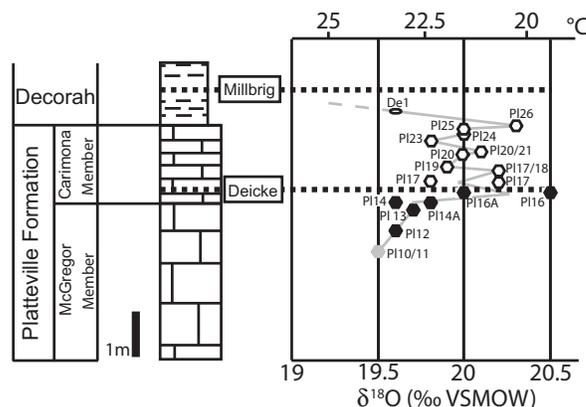


Figure 1. Plot of samples PI10/11 through De1 from Buggisch et al. (2010). PI samples are from the Cummingsville Annex Section; stratigraphy is from Webers (1966). Note that, using the correct placement of samples with regard to Deicke K-bentonite (DKb) bed, the cooling event occurred prior to the volcanic eruption. Black solid polygons are samples from below the DKb, black open polygons are samples above the DKb within the Carimona; oval is the first sample within the Decorah. Lowermost sample (gray polygon) reported in the supplementary data set of Buggisch et al. (2010, GSA Data Repository item 2010087) is PI10/11, but it apparently was not plotted in their figures. Additionally, this sample (PI10/11) is listed as 3.50 m in the supplementary data set, but according to the Webers (1966) stratigraphy, is at ~6.46 m.

words, the answer to the question "Did intense volcanism trigger the first Late Ordovician icehouse?," using the data set of Buggisch et al. (2010), seems to be "no." We have found the same pattern of apparent temperature change, i.e., cooling prior to the DKb, in the Sogn, Minnesota, section (Herrmann et al., 2010). Clearly, geographic coverage is still limited, but this agreement across studies argues that conodont paleothermometry will increasingly provide a valuable perspective on Paleozoic climate evolution.

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