Markov chain Monte Carlo based on adaptive Metropolis algorithm applied in combined model to deal with the uncertainty of weights of single models

Zhenxiang Xing, Han Zhang, Yi Ji, Gong Xinglong, Qiang Fu and Heng Li

ABSTRACT

Reliability and validity of model prediction play a decisive role in water resource simulation and prediction. Among many prediction models, the combined model (CM) is widely used because it can combine the prediction results of multiple single models and make full use of the information provided by various methods. CM is an effective method to improve the predictive veracity but the weight of single model estimation is the key to the CM. Previous studies take errors as the objective function to calculate the weight, and the uncertainty of the weight of the individual model cannot be considered comprehensively. In order to consider the uncertainty of the weight and to improve universal applicability of the CM, in this paper, the authors intend the Markov chain Monte Carlo based on adaptive Metropolis algorithm (AM-MCMC) to solve the weight of a single model in the CM, and obtain the probability distribution of the weight and the joint probability density of all the weight. Finally, the optimal weight combination is obtained. In order to test the validity of the established model, the author put it into the prediction of monthly groundwater level. The two single models in the CM are time series analysis model (TSAM) and grey model (GM (1,1)), respectively. The case study showed that the uncertainty characteristic of the weight in the CM can be obtained by AM-MCMC. According to the study results, CM has obtained a least average root mean square error (RMSE) of 0.85, a mean absolute percentage error (MAPE) of 8.61, and a coefficient of determination ($R^2$) value of 0.97 for the studied forecast period.

Key words | AM-MCMC, combined model, GM (1, 1), groundwater level, time series analysis

INTRODUCTION

Model construction is not only a key technology of water resources simulation and prediction, but also an important link with water resources management and sustainable utilization. The traditional approach to forecasting involves choosing the single forecasting method judged most appropriate of the available methods (Chambers et al. 1971) and applying it in some specific situations. The main methods of water resource simulation and prediction are time series analysis (Tingsanchali & Gautam 2015; Dastorani et al. 2016) and system analysis (Nayak et al. 2004; Dejamkhooy et al. 2017). In general, the choice of a method depends upon the characteristics of the series and the type of application (Makridakis et al. 1985). Therefore, using different prediction methods will achieve different prediction results. If the prediction error is simply used as the evaluation standard, some useful information will be lost (Guo et al. 2009). For example, some models like grey model (GM (1,1)) are not suitable for solving the periodic time series (Tseng et al. 2003), and other models like time series analysis model (TSAM) are poor to deal with the grey problem with insufficient data, small sample, complete information and lack of experiences (Chang & Tsai 2008). Hence the
full expression of the series of features cannot be achieved by solely relying on a single forecasting method. It is the reason why the combined model (CM) was put forward. The CM that combines the different forecasting methods appropriately was first proposed by Bates & Granger (1969), which eliminates the problem that individuals have no choice but to use a single method and rely exclusively on its forecasts. Whereas there is a key problem in the CM that it is difficult to estimate the weight coefficient of the individual model (Liu et al. 2013). CM based on entropy weight (EW-CM) and set pair analysis (SPA-CM) was a common method to solve the weight of single model in the CM (Yuan et al. 2014). The core of common method is to construct the objective function by means of absolute error or relative error, and the objective function is minimized under certain criteria. However, the simple arithmetic average method cannot accurately express the uncertainty of the weight of a single model. Moreover, the prediction results of each model, whether a single model or synthetic model, can provide only the deterministic prediction value of the hydrological process, and the uncertainty of the prediction results cannot be quantitatively described; thus, the accurate prediction cannot be obtained (Liang et al. 2015).

Solution of the weights used in combining forecasts can be based on posterior model probabilities within a Bayesian framework. This procedure, which is typically referred to as Bayesian model averaging (BMA), is a standard approach to model uncertainty within the Bayesian paradigm, where it is natural to reflect uncertainty through probability (Steel 2011). Therefore, by applying the BMA method to obtain the weight of the CM, the uncertainty of the weight of the individual model can be considered (Sloughter et al. 2007). There are two methods in BMA (Tian et al. 2012): expected to maximize (EM) and the Markov chain Monte Carlo algorithm (MCMC). The MCMC method can overcome the defect of the EM method (Gibbons et al. 2008), especially for two-parameter expressions, and the global optimal solution can be obtained visually. The convergence rate of the MCMC based on the adaptive Metropolis algorithm (AM-MCMC) was the fastest with the same accuracy (Marshall et al. 2004) compared with other sampling algorithms.

Given the uncertainty of the weight of a single model in the CM, a TSAM (Dastorani et al. 2016) and grey model (GM (1, 1)) (Dejamkhooy et al. 2017) are selected as the single models in the study problems in this work, which vary from simple to complex. In addition, the input and output are the same, but the basic theories are different for the TSAM and GM (1, 1) (see the Supplementary file, available with the online version of this paper). Under the theoretical framework of BMA, the AM-MCMC was applied to obtain the weight of the CM to construct the AM-MCMC-CM. Not only the optimal weight combination can be obtained, but the uncertainty of the weight of the individual model is taken into consideration. To verify the validity of the AM-MCMC-CM, it is applied to predict the groundwater level of Daxing Farm on the Sanjiang Plain in Heilongjiang Province, Northeast China. The results are compared with the statistics from the SPA-CM and another two single models.

DATA SET AND STUDY AREA

The Sanjiang Plain, an alluvial flat formed by the Heilongjiang River, Songhua River and Wusulijiang River, is located in northeastern Heilongjiang Province, China. The area of the Sanjiang Plain is 108,900 km². The mean annual volume of groundwater resources is 85.56 × 10⁸ m³. Recently, paddy planting has developed rapidly, and groundwater utilization for irrigation has increased annually. The groundwater level of the Sanjiang Plain has considerably changed. Therefore, it is important to analyze and forecast groundwater levels, which can provide references for decision making and groundwater utilization in the Sanjiang Plain Irrigation Area. Daxing Farm (132°18′ E–133°18′ E, 46°49′ N–47°13′ N), one of 14 large irrigation areas that encompasses 800 km², was selected as a typical study area in the irrigation region of the Sanjiang Plain. The geographical location of the study area is shown in Figure 1(a). The average monthly groundwater level at this farm over 17 years (January 1997 to December 2013) was obtained based on data collected at three groundwater monitoring wells (A (133°02′14.68″ E, 46°59′14.09″ N), B (133°10′44.7″ E, 47°07′44.75″ N), and C (132°54′41.39″ E, 47°04′26.48″ N)) by the Water Affairs Bureau of Heilongjiang Province. As shown in Figure 1(b), the time series variation of the average monthly groundwater level was given.
Establishment of a new CM

In this paper, the AM-MCMC-CM steps are as follows (Haario 2001):

1. First, initialize \( t = 0 \).
2. Randomly produce an initial state \((w_1, w_2)\) within \([0, 1] \) and use it as follows:
   (a) Calculate the covariance \( C_t \) according to Equation (1):
   \[
   C_t = \begin{cases} 
   C_0 & t \leq t_0 \\
   s_d(Cov(X_0, X_1, \cdots, X_{t-1}) + s_d \epsilon I_d) & t > t_0 
   \end{cases}
   \]  
   where \( C_0 \) is a diagonal matrix and is equal to 10% of the length of the search range; \( s_d \) is a scaling parameter that depends on only dimension \( d \), ensuring an acceptance probability in a reasonable range; \( X_0, X_1, \cdots, X_{t-1} \) is the sample that has been sampled; \( \epsilon \) is a constant; and \( I_d \) denotes a \( d \)-dimensional identity matrix. Gelman et al. (1996) suggested that \( s_d \) equals \( 2.4^2/d \).
(b) Produce a candidate state \( w_1^* \) and \( w_2^* \).
(c) Calculate an acceptance probability \( \alpha \) according to Equation (2):
   \[
   \alpha = \min\left(1, \frac{\sigma_{w_1^*, w_2^*}^2}{\sigma_{w_1^*, w_2^*}^2} \right)
   \]  
   where \( \sigma \) is the variance of the residual of the combined prediction value and the observed value.
(d) Generate a uniform random number \( u \) from \((0, 1)\).
(e) If \( u < \alpha \), then accept \( w_1^* \) and \( w_2^* \); otherwise, reject \( w_1^* \) and \( w_2^* \).
3. For \( t = t + 1 \), repeat (a)–(e) until the number of samples is sufficiently large.
4. According to the criterion of adaptive Metropolis convergence (Gelman & Rubin 1992), when AM-MCMC converges, the maximum value of the joint density \( f_{\max}(w_1^*, w_2^*) \) of \( w_1 \) and \( w_2 \) can be determined at the optimal point \((w_1^*, w_2^*)\) or the mean values of \( w_1^* \) and \( w_2^* \). The CM can be constructed as follows:
   \[
   H_{AMCM,t} = w_1^* \hat{H}_{TSAM,t} + w_2^* \hat{H}_{GM,t} 
   \]  
   (\( t = 1, 2, \cdots, n \))
   where \( H_{AMCM,t} \) is the output of the CM based on AM-MCMC, \( \hat{H}_{TSAM,t} \) is the output of an individual model (TSAM), and \( \hat{H}_{GM,t} \) is the output of the other individual model (GM (1, 1)).
The basic block diagram of the proposed prediction approach is shown in Figure 2.

Model performance evaluation

Three statistical indicators, the mean relative error (MAPE), root mean square error (RMSE), and coefficient of determination ($R^2$), were used to test the performance of the CM (Ramana et al. 2013; Soleymani et al. 2016).

RESULTS AND ANALYSIS

Stochastic analysis of weights

The weight set $(w_1, w_2)$ of the two individual models above was sampled using the AM-MCMC algorithm. The iteration was set to 6,000 in the initial period, and parallel sampling was performed five times. Twenty thousand samples were produced at one time; therefore, $(20,000-6,000) \times 5 = 70,000(w_1, w_2)$ sets were generated and used to study the weighting uncertainty of the CM. As shown in Figure 5(a), the value of $R^{1/2}$ decreases to less than 1.02 and steadily declines to 1 after the 6,000th iteration. This result suggests that the multiple sampling series converged, i.e. the statistical characteristics of the samples have converged the ensemble. As shown in Figure 3(b) and 3(c), the posterior marginal densities of the CM weights are $w_1 \sim N (0.6147, 0.21^2)$ and $w_2 \sim N (0.3853, 0.17^2)$. Meanwhile, the posterior joint probability densities of the two weight parameters are shown in Figure 3(d), and the weights reached the global maximum joint density at the intersection of the mean values $(w_1^*, w_2^*) = (0.6147, 0.3853)$; thus, $f(w_1^*, w_2^*)_{\text{max}} = 7.8$.

Calculation results based on AM-MCMC-CM

Through the above calculation, the model expression of the TSAM, GM (1, 1), AM-MCMC-CM, SPA-CM are given in Table 1. In Equations (4)–(7), $t$ represents the time series. Finally, the corresponding prediction value is obtained according to the model expression.

The scatter plots of observed values and the predictions of the four models in the forecast period (2011–2013) is shown in Figure 4. As shown in Figure 4, the calculated values of the four types of prediction models exhibit good correlation with the measured groundwater levels. Among the models, the prediction of TSAM and GM (1, 1) are quite similar, but the effects of the two single models are very different in low and high groundwater level prediction, meaning that the choice of the two single models is successful. Additionally, this also explains the necessity of using a CM. It can be seen that the predictive effect of the CM is clearly better than that of the single model from the scatter graph.

Model evaluation and discussion

The model evaluation indicators MAPE, RMSE and $R^2$ introduced in section ‘model performance evaluation’ were used to further compare the accuracies of the AM-MCMC-CM,
SPA-CM, TSAM and GM (1, 1) models. The evaluation indicator values of the four models in the calibration period (1997–2010) and forecast period (2011–2013) are listed in Table 2, and the predicted (dashed lines) and observed (solid lines) levels in the forecast period (2011–2013) are plotted in Figure 4.

As shown in Table 2, in the calibration period (1997–2010), the MAPE, RMSE and $R^2$ were 9.52%, 0.84, and 0.91, respectively, for TSAM, and 10.98%, 0.97, and 0.88, respectively, for GM (1, 1). Additionally, as shown in Figure 4, TSAM yielded a better fit with observations than did GM (1, 1), which implies that TSAM should be assigned a larger weight than GM (1, 1). This result is in accordance with the CM results based on Equations (6) and (7). Likewise, the results of the analysis of correlation and the dispersion degree in the previous section ‘Calculation results based on AM-MCMC-CM’ suggested that the three evaluation indicators of the CM are better than those of each individual model from 1997 to 2013, as shown in Table 2. Furthermore, the results indicate that the CM is more accurate than individual models in predicting the groundwater level.
Table 1 | The equations used in models of groundwater level prediction

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<th>Type</th>
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| TSAM (Dastorani et al. 2016) | \( \hat{H}_{TSAM} = -0.0001t^2 + 0.030478t + 5.1934 + 0.000066264 - 0.5922 \cos \frac{4\pi t}{168} - 0.1704 \sin \frac{4\pi t}{168} \) \( (4) \)  
  \(- 1.9434 \cos \frac{28\pi t}{168} - 0.4463 \sin \frac{28\pi t}{168} + 0.8627 \cos \frac{56\pi t}{168} + 0.2775 \sin \frac{56\pi t}{168} + 0.5784t_{t-1} \)  
  \(- 0.1715t_{t-4} - 0.2020t_{t-6} - 0.1458t_{t-10} + \varepsilon_t \in (0, 0.7016) \) \( t = 1, 2, \ldots, n \) |
| GM (1, 1) (Tseng et al. 2001) | \( \hat{H}_{GM,t+1} = (1 - e^{-0.002}) \left[ H^{(0)}(1) + \frac{71.338}{0.002} \right] e^{0.002t} \) \( t = 1, 2, \ldots, n \) \( (5) \) |
| AM-MCM-CM                     | \( \hat{H}_{AMCM,t} = 0.6147 \hat{H}_{TSAM,t} + 0.3853 \hat{H}_{GM,t} \) \( (6) \) |
| SPA-CM (Guo et al. 2009)      | \( \hat{H}_{SPA-CM,t} = 0.5629 \hat{H}_{TSAM,t} + 0.4371 \hat{H}_{GM,t} \) \( (7) \) |

Figure 4 | Scatter plots of observed and predicted groundwater levels at Daxing Farm from 2011–2013: (a) TSAM, (b) GM (1, 1), (c) SPA-CM, and (d) AM-MCM-CM.
The AM-MCMC-CM results for the calibration period yielded MAPE, RMSE and $R^2$ values of 8.24%, 0.76 and 0.98, respectively. Thus, AM-MCMC-CM produced better results for almost all the evaluation indicators considered in this study compared to those of the other models. Both AM-MCMC-CM and SPA-CM produced linear correlation coefficient (CC) values greater than 0.85 based on the associated scatter plots (Figure 4). In addition, a strong correlation exists between the observed and simulated groundwater levels of both models. However, a higher degree of correlation was obtained by AM-MCMC-CM than by SPA-CM (not considering the uncertainty of the weights of individual models).

Through statistical and graphical analysis and comparison of evaluation indicators, the prediction accuracy of the CM was better than that of individual models. AM-MCMC-CM performed better than SPA-CM for groundwater level prediction; notably, the RMSE was 4.0% lower in the calibration period (1997–2010) and 5% lower in the forecast period (2011–2013). In addition, the MAPE decreased by 0.76% and 0.70%, and $R^2$ increased by 2.0% and 2.1%, respectively. AM-MCMC-CM considers the weighting uncertainty of individual models to reduce the random interference of the weights of individual models caused by observed data limitations. Notably, the weight of TSAM, which was the most accurate individual model, was increased from 0.5629 in SPA-CM to 0.6147 in AM-MCMC-CM. Overall, AM-MCMC-CM exhibited higher reliability and accuracy than SPA-CM in groundwater level forecasting.

### CONCLUSIONS

The AM-MCMC-CM is based on the AM-MCMC algorithm to consider the uncertainty of the weight of a single model in the CM and applied to predict the groundwater level to verify the effectiveness of the new model. The conclusions are as follows:

1. The AM-MCMC-CM can account for the uncertainty and equi-finality of the different weights of individual models. Furthermore, it can obtain the joint probability distribution and statistical characteristics of weights of individual models used in the CM.
2. AM-MCMC-CM can effectively reduce the influence of the prediction error of individual models on the prediction results of the CM compared with that of the SPA-CM. Moreover, the proposed model weakens the dependence on a limited number of samples in the determination of weights.
3. AM-MCMC-CM is more applicable and provides better prediction accuracy than that of the TSAM, GM (1, 1) and SPA-CM. Thus, AM-MCMC-CM can improve the accuracy of groundwater level prediction.

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<td>AM-MCMC-CM</td>
<td>SPA-CM</td>
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<td>MAPE</td>
<td>8.24%</td>
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<tr>
<td>RMSE</td>
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